

RADAR-BASED NOWCASTING TECHNIQUES (WITH PRACTICALS)

WMO Severe Weather Forecasting Programme (SWFP) Regional Sub-programme for Southeast Asia (SWFP-SeA) Training Desk and Study Visit for Cambodia (Ha Noi, 19 - 23 May 2025)

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Weather Radar

• Radar - RAdio Detection And Ranging

- Weather radar emits microwave radiation into the atmosphere and receives the reflected radiation called <u>echoes</u> from water droplets in the air
 - Microwave in the solar radiation and the Earth's emission spectrum is not intense
 not affecting the operation of weather radar
 - Moving speed of radar echoes (towards or moving away from radar) can be detected based on Doppler effect







) WMO

Operation principle of weather radar (I)



Frequency	Band	Wavelength	Applications	
400 - 900 MHz	UHF	0.3 - 0.7 m	wind profiler	
I GHz	L-band	0.3 m	boundary layer wind profiler	
2 - 4 GHz	S-band	7-15 cm	long range precipitation radar	
4 - 8 GHz	C-band	4-7 cm	long range precipitation radar	
8 - 16 GHz	X-band	2-4 cm	precipitation radar	
16 - 20 GHz	Ku-band	I-2 cm	precipitation / cloud radar	
35 Hz	Ka-band	8.5 mm	precipitation / cloud radar	
90 - 100 GHz	W-band	3 mm	cloud radars	





Operation principle of weather radar (2)

- Each conical scan is called a Plan Position Indicators (PPI)
- Collection of all conical scans is called a volume scan
- Horizontal cross-section is called Constant Altitude PPI (CAPPI)



• Vertical cross-section is called Range Height Indicator (RHI)







ΜΟ

Operation principle of weather (3)

Most weather radars are pulsed radars, i.e. sending out pulses of microwave.



Time between pulses = Pulse repetition time PRT or Pulse repetition interval PRI (T)

Pulse repetition frequency (PRF) = I/T

- Pulse width τ is usually ~ micro-second (10⁻⁶ s)
- Longer pulse width means more energy being transmitted, can improve detection of weaker echoes at longer range.
- Longer pulse limits the spatial resolution of radar data
- PRF (inverse of PRT) is usually 100 3,000 Hz
- Smaller PRF (longer PRT) enables detection of echoes to greater distance (before emission of next pulse)





Operation principle of weather radar (4)

The received power (P_R) at the radar of the returned microwave due to scattering of raindrops can be derived as the following <u>radar equation</u>:



Introduce a new parameter reflectivity, which is a logarithmic parameter measured in unit of dBZ:

$$Z = 10\log_{10}\left(\frac{z}{1\,mm^6\,/\,m^3}\right)$$

Reflectivity factor z is related to the rainfall rate R by an empirical relation, z-R relation:

 $z = a R^b$ z in unit of mm⁶/m³; R in unit of mm/h

For precipitation due to stratiform clouds 層狀雲, the relation is known as the Marshall-Palmer relation:

$$z = 200 R^{1.6}$$

➔ Accurate estimation of coefficient values in z-R relationship is important for monitoring and forecasting rainfall from weather radars





Reflectivity and Rainfall Rate

Reflectivity factor (Z) measured by weather radar is the sixth moment of drop size distribution N(D) (D=drop diameter in mm):

$$Z = \int_0^\infty N(D) D^6 dD$$

Rain rate (R) is given by:

$$R = \frac{\pi}{6} \int_0^\infty N(D) D^3 V(D) dD$$

V(D) is the fall velocity of particles of size D

Z = Z (R, N(D))

an empirical relationship on Z-R:

 $z = a R^b$



$$Z = aR^{b}$$
$$dBZ_{i} = IOlog(a) + b dBR_{i}$$

$$dBZ = 10 \log_{10} Z$$

 $dBR = 10 \log_{10}R$





From Radar Reflectivity to Rainfall



With the Z-R relation, can convert dBZ \rightarrow dBR \rightarrow hourly rainfall

Scatterplots of estimated radar rainfall vs actual

 $z = a R^b$

dynamical update a and b

static a and b based on historical data







Converting Radar Reflectivity Image to Rainfall Rate

In weather radar image of reflectivity, different reflectivity (in dBZ) is assigned different colours for identifying where heavy rain is located.

 $dBZ = 10 \log_{10} Z$

Level	dBZ	Description	Rainfall rate in mm/h	Rainfall rate in mm/h	Rainfall rate in mm/h
			(Z = 200 R ^{1.6})	(Z = 300 R ^{1.4})	(Z = 58.8 R ^{1.56})
1	18-30	Light rain			
2	30-41	Moderate rain			
3	41-46	Heavy rain			
4	46-50	Very heavy rain			
5	50-57	Intense rain			
6	>57	Extreme rain			





Practical I



- I. Estimate the maximum rainfall rate (mm/hr) to the west of the Pearl River Estuary (area A)
- 2. Do you expect the actual I hour accumulated rainfall is similar, larger / smaller than the results in Q1? Explain your answer.





Comparing reflectivityrainrate (Z-R) relationships





Z-R in other Meteorological Services



DECIBEL OF RAINFALL RATE (DBR)

Gridded QPE from Radar and Rain Gauge Observations

KONG OBSERVATORY

Barnes Analysis

G

G

- interpolation with Gaussian weighting according to distance between data & estimation point
- consider correction using residuals and grouping of rain gauges

$$B(x_0) = \frac{\sum_{i=1}^{N_0} w_i G_i}{\sum_{i=1}^{N_0} w_i}$$

$$B : \text{Barnes estimation (mm)}$$

$$L : \text{radius of influence}$$

$$N_0 : \text{number of gauge report}$$

$$G : \text{i-th gauge report (mm)}$$

$$w_i : \text{weight of i-th gauge}$$

$$h_i : \text{distance between gauge and estimation point}$$

Co-kriging Analysis

- extends from ordinary geostatistical kridging technique and take into account both radar reflectivity and rain gauge data
- consider error characteristics (spatial covariance and correlation) of rain gauge measurements and radar reflectivity field



$$\sum_{i=1}^{N_0} \lambda_i(x_0) \gamma_{GG}(x_n, x_i) + \sum_{j=1}^{M_0} \lambda_j(x_0) \gamma_{GR}(x_n, x_j) + \mu_G(x_0) = \gamma_{GG}(x_n, x_0), \quad \text{for } n = 1, \dots, N_0$$
$$\sum_{i=1}^{N_0} \lambda_i(x_0) \gamma_{RG}(x_m, x_i) + \sum_{j=1}^{M_0} \lambda_j(x_0) \gamma_{RR}(x_m, x_j) + \mu_R(x_0) = \gamma_{RG}(x_m, x_0), \quad \text{for } m = 1, \dots, M_0$$



Example of Barnes vs Co-kriging methods







Dual-polarization Radar

Dual-polarization radar transmit and receive both horizontal and vertical polarized pulses. •



Raw Data

- Reflectivity (Z)
- Velocity (V)
- Spectral Width (W)
- Differential Reflectivity (ZDR)
- Differential Phase (Φ_{DP}) ٠
- Specific Differential Phase (K_{DP})

Weather Forecasting Programme Regional Sub-programme for Southeast Asia (SWFP-SeA)

Correlation Coefficient (ρ_{hy})

Differential Reflectivity (ZDR)

Useful indicator of the average drop shape of the ٠ dominant hydrometeor within the resolution volume



Physical Interpretation of ZDR	Major Axis Diameter (mm)
≻Rain →	< 0.3 mm
≻Hail tends to tumble	1.35 mm
 appears spherical to radar ZDR ~ 0 dB 	1.75 mm
	2.65 mm
Small and completely water-coated hail	2.90 mm
 appears as giant raindrop ZDB ~ 5 - 6 dB 	3.68 mm
	4.00 mm

Non-met echoes

- Ground clutter: very noisy
- Biological scatterers: less noisy, orientation with flight direction

\rightarrow Dual-pol parameters can improve identification of rain drops and types of hydrometeors, and estimation of rainfall from radar

ZDR (dB)

~ 0.0 dB

~ 1.3 dB

~1.9 dB

~2.8 dB

~3.3 dB

~4.1 dB

~4.5 dB

Data from Wakimoto and Bringi (1988)

Image

0

0

0

0



Doppler Velocity



- Noted that the Doppler frequency shift is rather small compared to the transmitted frequency of microwave (why?)
- Doppler radars detect air motion by measuring phase shift of microwaves







Doppler Velocity



Given a known wind vector (\mathbf{M}) at a point, only the radial component of the total wind (M_r) is detected by Doppler weather radar.

- At point (a) the total wind vector \mathbf{M} is along the radial line, hence the radar can detect the whole wind. ٠
- At point (b) the total wind vector **M** is tangent to the circle, hence the radar sees **zero** wind there. ٠
- At point (c) the total wind vector \mathbf{M} has a radial component M_r that the radar detects, and a tangential ٠ component M_{tan} that is invisible. Thus only part of the total wind is detected there.





Ouestion: How could the true wind field be determined using radar?

Oppositely if the full wind vector is unknown. Beware that an infinite number of true wind vectors can create the same radial component as detected by the Doppler radar.

- For example, at point (d) two completely different wind vectors M_1 and M_2 have the <u>same</u> radial component M_r
- The radar only detects M_r but there is insufficient information to determine the total wind vector (speed) unless under certain specific condition such as the weather system or associated features taking direction of movement in radial direction towards / away from the radar





Examples of wind flow or condition for Doppler radial velocity field (warm color +ve / radial outward; cold color –ve / radial inward);

Radar is located at the centre of the circles.





Convergence / Divergence

Look for 'dipoles' of warm/cold colors <u>along a radial</u> to identify convergence/divergence





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Cyclonic / Anti-cyclonic Vortex

Look for 'dipoles' of warm/cold colors on <u>either side of a</u> <u>radial</u> to identify cyclonic/anti-cyclonic signatures



Cyclonic signature (anti-clockwise)



Anti-cyclonic signature

Radar

(clockwise)



Typhoon Higos (19 Aug 2020)

TMS 0 deg PPI Reflectivity

TMS 0 deg PPI Doppler Velocity







- Thunderstorms are usually associated with intense echoes (e.g. greater than 50 mm/hr)
- Lightning location images are helpful to monitor thunderstorms



3 KM CAPPI Rainfall Rate

Lightning location images overlaid on CAPPI Reflectivity





🆄 wmo



credit: HKO



-5/



15

10

NWG

1

nfall

10

0.50

0.15

Bow-shaped echoes in a squall line signifying severe convective weather such as violent wind gust



Scale: 20 to 200 km

The bow-shaped echo is a result of focusing of the strong flow at the rear of the system.





Bow-shaped echoes in a squall line signifying potential damaging winds





Storm	34	47
Gale	24	34
StrongF7	20	24
StrongF6	16	20
Fresh	11	16
Moderate	4	11
Light	1	4
Calm	-1	1
Light	-4	-1
Moderate	-11	-4
Fresh	-16	-11
StrongF6	-20	-16
StrongF7	-24	-20
Gale	-34	-24
Storm	-47	-34

Beaufort Scale_0.7 Colour Scheme

Case II: 2016-04-12

- Identify the direction of wind flow and location of shear line from the radar images below
- Estimate the maximum wind (gust) recorded when the rain band pass over the vicinity of the HK International Airport (Point A)







Can you identify any hazardous convective weather system from the following radar images?









Nowcasting of Precipitation and Thunderstorms using Radar Image

- Determine the motion of radar echoes from successive radar images (e.g. every 6 minutes)
- <u>Assume the intensity of echoes does not change</u>, extrapolate radar echoes along the motion vectors and determine the future location of radar echoes in next I-6 hours
- Derive forecast rainfall from forecast echo location

 intensity (reflectivity) of radar echoes depending
 on size and number density of rain drops
- Limitation in radar extrapolation nowcast
 - no change in intensity
 - motion of echoes remains the same





QPF - tracking motion of radar echoes

Maximum Correlation



where Z_1 and Z_2 are the reflectivity at T+0 and T-6min respectively

$$\mathbf{R} = \frac{\sum_{k} Z_{1}(k) \times Z_{2}(k) - \frac{1}{N} \sum_{k} Z_{1}(k) \sum_{k} Z_{2}(k)}{\left[\left(\sum_{k} Z_{1}^{2}(k) - N \overline{Z_{1}}^{2} \right) \times \left(\sum_{k} Z_{2}^{2}(k) - N \overline{Z_{2}}^{2} \right) \right]^{1/2}}$$

Optical Flow

Given I(x,y,t) the image brightness at point (x,y) at time t and the brightness is constant when pattern moves, the echo motion components u(x,y) and v(x,y) can be retrieved via <u>minimization of the cost function J</u>:

$$J = \iiint \left[\frac{\partial I}{\partial t} + u\frac{\partial I}{\partial x} + v\frac{\partial I}{\partial y}\right]^2 dxdydt$$



I-hr nowcast reflectivity using optical flow versus TREC motion fields





Given the following sequence of radar image, estimate:

(a) Moving speed of intense rainband (i.e. front-edge of yellowish radar echo areas) over the western coastal areas of Guangdong

(b) When would it start to affect Hong Kong?





























2025-03-10

10:30 UTC



When would it expect to affect Phnom Penh?

11.41N 104.89E

2025-03-10

11:00 UTC



When would it expect to affect Phnom Penh?

11.41N 104.89E

2025-03-10

11:30 UTC



When would it expect to affect Phnom Penh?

11.41N 104.89E





- Fitting elliptical object based on selected reflectivity thresholds • on single and multi-layer CAPPI.
- Determine intensity weighted centroid (centre of ellipse) and ٠ geometric (principal component) parameters including major axis, minor axis and orientation angle (θ)
- Specify the storm object properties for tracking and identifying • severe weather characteristics: hail signature, gust, lightning and rainfall intensities



62









High-impact hazards of hail, violent gust / convective downburst, significant lightning and heavy rain within the next 30 minutes in Hong Kong area (red box)





SWIRLS Downburst Nowcast Algorithm







Indicators of the cloud-to-ground (C-G) lightning onset

- effective indicators in predicting onset of C-G lightning in 10-20 minutes:
 - TOP > 7 km
 - VIL > 5 mm
 - $\text{REF}_{0^{\circ}\text{C}} > 42 \text{ dBZ}$
 - $\text{REF}_{-10^{\circ}\text{C}} > 17 \text{ dBZ}$

 $I = -3.623 \times 10^{-01}$ $+6.105 \times 10^{-02} \times TOP(km)$ $+2.601 \times 10^{-02} \times VIL(mm)$ Lightning Intensity $+1.967 \times 10^{-06} \times REF_{-20^{\circ}C}(Z)$ $+1.146 \times 10^{-07} \times REF_{0^{\circ}C}(Z)$ + • Observ $flash_rate = 10^{I} = 10^{a+\sum b_i x_i}$ $I = a + \sum b_i x_i$

Predicted Lightning Intensity

 $- \text{REF}_{-20^{\circ}\text{C}} > 0 \text{ dBZ}$



Hail Detection

- Hail
 - Large reflectivity (> 60 dBZ) above freezing level with weak echo region (WER) underneath
 - Overhang signature in vertical cross section
- Indicator
 - Echo top (TOPS) of 58-dBZ > 3 km
 - Vertical Integrated Liquid (VIL) in 0-2 km layer < 5 mm



Severe Thunderstorms



- Updraft is strong, with sign of rotation (mesocyclone).
- In the strongest updraft, the air becomes supercooled even at temperatures as low as -40°C. On weather
 radar, this region of supercooled updraft air appears as weak echo region. This is a characteristics for hail
 formation.
- Accompanied with weak echo region, there is often overhang, which is another characteristics for hail formation.











Bounded Weak Echo Region



in the PPI; a "grim reaper" shape representing the BWER in the RHI.





"The presence of a BWER indicates the thunderstorm possesses **an intense updraft** ... BWERs are almost always associated with **supercells** ...

Persistent BWERs, as reasonably reliable proxies for supercells, suggest that the storm is capable of producing **damaging winds, large hail** and, on occasion, tornadoes...." (meted: radar for severe convective weather)

Reddish top up to 10km Intense core (> 60dBz)around 5-6 km





WMO Seve Trair

Deep Learning Nowcast







AI in Precipitation Nowcasting



Deep Learning Model



Two-hour radar reflectivity nowcasts from optical flow and DL models







- Better at predicting the movement and intensity evolution radar reflectivity than extrapolation using gridded optical flow field
- Generally, sequence-to-sequence video prediction model trained with simple loss function (MSE / MAE) faces the blurring problem.



Input

sequence

Input

sequence

ResConvLSTM-GAN

- ConvLSTM with residual connections in encoderforecaster network
- Generative Adversarial Network (GAN) to improve representation of small-scale features







ResConvLSTM-GAN



Application of ResConvLSTM-GAN





Evaluation metrics (I)

- Accuracy measures
 - Probability of Detection (POD)
 - False Alarm Ratio (FAR)
- Skill measures
 - Critical Success Index (CSI)
 - Note CSI is not equitable score as CSI is nonzero for constant "yes" forecast or random forecast

$$CSI = \frac{H}{H + M + F}$$

 $POD = \frac{H}{H+M}$

 $FAR = \frac{F}{H+F}$

Observed Yes No False Hit Yes Alarm (H/TP)Forecast (F / FP) True Miss Å Negative (M / FN) (Z / TN)

- Heidke Skill Score (HSS)
 - An equitable score as HSS is zero for constant or random forecast

$$\begin{split} HSS &= \frac{POC - POC_{\texttt{random}}}{1 - POC_{\texttt{random}}} = \frac{(H + Z) - (H_{\texttt{random}} + Z_{\texttt{random}})}{N - (H_{\texttt{random}} + Z_{\texttt{random}})} \\ \\ H_{\texttt{random}} &= \frac{(H + M)(H + F)}{N} \qquad Z_{\texttt{random}} = \frac{(Z + M)(Z + F)}{N} \end{split}$$

N = no. of samples = H+F+M+Z

Percentage of correct (POC):

$$POC = \frac{H}{H+Z}$$



Evaluation Metrics (2)

- Fraction Skill Scores (FSS)
 - Fuzzy verification to better represent forecast skills due to double penalty

$$FSS = 1 - \frac{\sum_{N} (P_{fcst} - P_{obs})^2}{\sum_{N} P_{fcst}^2 + \sum_{N} P_{obs}^2}$$







Evaluation Metrics (3)

- Learned Perceptual Patch Similarity (LPIPS)
 - also known as "Perceptual Loss"
 - a neural network (NN) based metric to match human perception → a lower LPIPS score indicates a higher similarity
 - measure the difference of activations in a pre-defined NN between 2 images
 - a common pre-defined NN: AlexNet







• AI Methods > ROVER;



RCLG produces less blurry output, so it has lower POD and FAR





rover_nonlinear
 trajGRU
 ResConvLSTM-GAN

• RCLG has better overall skills compared to TrajGRU

Verification







LPIPS - Lead time plot



Verification

- Learned Perceptual Patch Similarity (LPIPS) metric
 - A Neural Network(NN)-based metric to match human perception, lower the better
 - Measure the difference of activations in a predefined NN between 2 images
 - Pre-defined NN: AlexNet
 - TrajGRU-generated outputs are perceptually dissimilar while RCLG could generate more realistic outputs

Recent Updates on Deep Learning of Radar Nowcasting







T+120 min nowcasts from deep learning models

2025-05-07 II:18 HKT









Collaboration with **PAGASA** and **TMD** in knowledge transfer and implementing Com-SWIRLS and deep learning nowcast models











weather.tmd.go.th/composite/index_qpe.html 🞛 🛛 🥨 Cisco Webex Meeti... 🏾 🏄 Data Science Bootc... 🛛 🖬 Monkeytype 🛛 🥐 ChatPDF - Chat

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19.24°

17.24°N

SNK

103.87°E

Deep learning for precipitation nowcasting: A benchmark and a new mode

香港天文台

HONG KONG OBSERVATORY

105.87°E

KKN

101.87°E



Thank you very much

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17120°C 80-95%

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